# Carbon pool in soil under organic and conventional farming systems

Magdalena Hábová<sup>1</sup>, Lubica Pospíšilová<sup>1</sup>, Petr Hlavinka<sup>2,3</sup>, Miroslav Trnka<sup>2,3</sup>, Gabriela Barančíková<sup>4</sup>, Zuzana Tarasovičová<sup>4</sup>, Jozef Takáč<sup>4</sup>, Štefan Koco<sup>4</sup>, Ladislav Menšík<sup>5</sup>, Pavel Nerušil<sup>5</sup>

**Citation**: Hábová M., Pospíšilová L., Hlavinka P., Trnka M., Barančíková G., Tarasovičová Z., Takáč J., Koco Š., Menšík L., Nerušil P. (2019): Carbon pool in soil under organic and conventional farming systems. Soil & Water Res., 14: 145–152.

Abstract: Changes in the agricultural management and climatic changes within the past 25 years have had a serious impact on soil organic matter content and contribute to different carbon storage in the soil. Prediction of soil carbon pool, validation, and quantification of different models is important for sustainable agriculture in the future and for this purpose a long-term monitoring data set is required. RothC-26.3 model was applied for carbon stock simulation within two different climatic scenarios (hot-dry with rapid temperature increasing and warm-dry with less rapid temperature increasing). Ten years experimental data set have been received from conventional and organic farming of experimental plots of Mendel University School Enterprise (locality Vatín, Czech-Moravian Highland). Average annual temperature in this area is 6.9°C, average annual precipitation 621 mm, and altitude 530 m above sea level. Soil was classified as Eutric Cambisol, sandy loam textured, with middle organic carbon content. Its cumulative potential was assessed as high. Results showed linear correlation between carbon stock and climatic scenario, and mostly temperature and type of soil management has influenced carbon stock. In spite of lower organic carbon inputs under organic farming this was less depending on climatic changes. Conventional farming showed higher carbon stock during decades 2000–2100 because of higher carbon input. Besides conventional farming was more affected by temperature.

Keywords: crop management and climatic scenarios; RothC-26.3 model; soil organic carbon

Current status and changes in soil organic carbon stock as response to agronomic and climatic conditions become extremely important today. There is an effective strategy to mitigate global climate change by increasing carbon stock in soil (SMITH *et al.* 2010; MACHMULLER *et al.* 2015). The level and balance of soil organic carbon is also the main criterion of agricultural sustainability. The last depends

Supported by the National Agricultural Agency, Project Earth QK1810233, by the Ministry of Agriculture of the Czech Republic, Project No. RO0418, by the Slovak Research and Development Agency (APVV-0243-11 and APVV-14-0087), and by the Project SustES "Adaptation strategies for sustainable ecosystem services and food security under adverse environmental conditions" (CZ.02.1.01/0.0/0.0/16\_019/0000797).

<sup>&</sup>lt;sup>1</sup>Department of Agrochemistry, Soil Science, Microbiology and Plant Nutrition, Faculty of AgriScience, Mendel University in Brno, Brno, Czech Republic

<sup>&</sup>lt;sup>2</sup>Department of Agro systems and Bioclimatology, Faculty of AgriScience, Mendel University in Brno, Brno, Czech Republic

<sup>&</sup>lt;sup>3</sup>Global Change Research Institute, Czech Academy of Sciences, Brno, Czech Republic

<sup>&</sup>lt;sup>4</sup>National Agricultural and Food Centre, Soil Science and Conservation Research Institute, Prešov, Slovakia

<sup>&</sup>lt;sup>5</sup>Division of Crop Management Systems, Crop Research Institute, Prague-Ruzyně, Czech Republic

on soil ability to maintain productive and other non-productive functions (biodiversity provision, hygienic, environmental etc.). In this way soil organic carbon is regarded as a key factor influencing both of them. Whether soils are a sink or source of carbon depends on the current organic carbon stock, agricultural practices over time, soil properties (e.g. clay content, soil depth, content and quality of plant and organic input, fertilizing etc.), and climatic conditions (Baldock & Skjemstad 1999; Song et al. 2014). As quoted DE LIU et al. (2016) the amount of soil organic carbon that is attained under agriculture largely depends upon the carbon input and its decomposition rate under various agronomical practices. Today the agricultural measures encouraged soil conversion to the organic farming and minimum tillage technology, with aim to increase carbon stock in soil (SMITH et al. 2007; KACZYNSKI et al. 2013). On the other hand, conventional and intensive farming, simplification of crop rotation cause the decreasing of carbon stock. Sustainable soil management systems require the proper choice of crop rotation system, agricultural technics, carbon stock, as well as a supply of nutrients to reach the higher productivity (KING et al. 2005; LAMAR et al. 2006). LORENZ and LAL (2005) stress that conventional analytical methods for measuring of total organic carbon (TOC) are expensive, timeconsuming, and not always comparable. VISCARRA ROSSEL et al. (2016) demonstrate using of spectroscopic and gamma attenuation sensors for TOC stocks estimating. For their validation and quantification long-term monitoring data set is required. Widely used models for carbon stock prediction are RothC 26.3, CENTURY, CANDY, and DAISY. They were validated in Europe for the period of 1990-2080 (FALLOON et al. 1998, 2000; Ронанкоvá et al. 2015). RothC model was originally developed and parametrized to model turnover of organic carbon in arable soil from Rothamsted long-term field experiments. Later it was extended to model turnover in grassland and woodland and operates in different soils and under different climates (COLEMAN et al. 1997; SMITH et al. 1997, 2005, 2007; KERYN & POLGLAS 2004). It has also been set from an empirically-derived relationship between inert and total soil organic carbon content (SOC) (FALLOON et al. 1998, 2000; FALLOON & SMITH 2002). Inert organic carbon was according to Jenkinson et al. (1987, 1999) defined as a fraction of soil organic matter that is biologically inert and has an equivalent radiocarbon age of more than 50 000 years. Besides inert organic carbon, total

organic carbon includes relatively stable and labile carbon forms. Stable carbon forms are represented by carbon of humic acids, fulvic acids and humins (Stevenson 1994; Kučerík *et al.* 2007; Song *et al.* 2014). Humic substances can remain stored in the geosphere for thousands of years. Labile carbon forms are important from point of view soil biological activity. All of the organic carbon forms in soils are still not well studied and understanding of carbon sequestration is very important for evaluation of the global carbon cycle.

The aim of this study is to predict carbon sequestration under two different climatic scenario and crop management systems. Furthermore validation of RothC model for Cambisols, the most spread soil in the Czech Republic, is presented. Among the evaluated criteria are both quantitative and qualitative criteria of soil organic carbon.

## MATERIAL AND METHODS

Field experiments have been continuously conducted at locality Vatín (Czech-Moravian Highland). This area belongs to the potatoes growing area with average annual temperature 6.9°C, average annual precipitation 621 mm, and altitude 530 m a.s.l. Original Sanguisorba-Festucetum comutatae grassland (native) was ploughed and two crop sequences were chosen - organic and intensive crop sequences. Organic crop sequence (OCS) was represented by 33.4% of cereals, 16.6% of root crops, 16.6% of technical crops, and 33.4% of fodder. Nutrients were applied according to ratios (N-P-K, kg/ha/year), and involved 90-30-80 to winter wheat and 40-30-60 to spring barley; however 60% of inputs were in the organic form utilizing farmyard manure. Intensive crop sequence (ICS) was characteristic by more intensive agriculture and an optimal level of chemical inputs (mineral fertilizers, pesticides), but without organic farmyard manure. It was represented by 50% of cereals, 16.6% of root crops, and 33.4% of technical crops. Nutrients were applied at ratios (N-P-K, kg/ ha/year): 130-40-80 (winter wheat) and 60-35-80 (spring barley). A split plot method was used. Soil was sampled in the upper 0–20 cm Ap horizon twice a year (spring and autumn) during the period 1999-2016. The coordinates of soil profile were measured by Garmin Dakota 10 (Garmin International, Inc., USA) and are as follows: 49°31.091'N, 15°58.196'EO. Soil was classified according to the IUSS Working Group WRB (2015) as Eutric Cambisol. Horizon designation was done

by Jahn et al. (2006). Basic soil characteristics were determined by commonly used standard methods (ZBÍRAL et al. 2010). Soil reaction was determined by potentiometric method in distilled water and in 1M KCl solution (1:2.5). Particle size analysis was determined by the pipette method. Total organic carbon content was determined by oxidimetric titration method (Nelson & Sommers 1996). Fractional composition of humic substances was measured according to Kononova and Beltchikova method (1963, in: Pospíšilová et al. 2016). RothC-26.3 mode was set from an empirically-derived relationship between inert organic matter and total stock of organic carbon (FALLOON & Smith 2002; COLEMAN & JENKINSON 2005). Four active organic carbon forms in soil (decomposable plant material = DPM, resistant plant material = RPM, microbial biomass = BIO, humified organic matter = HUM), and inert organic carbon (IOC) were recognized. The incoming plant carbon is split between DPM and RPM, depending on their ratio. The decomposition rate is modified as a function of temperature, moisture and soil cover. The main model's input data are as follows:

Climatic data – monthly rainfall (mm), monthly evapotranspiration (mm), monthly air temperature (°C), Soil data – clay content (%), inert organic carbon content (%), initial organic carbon stock (t/ha), soil depth (cm),

Land use and management data – soil cover, monthly input of plant residues (t/ha), monthly input of organic manure (t/ha), residue quality factor (DPM/RPM ratio).

Climatic data were received from Meteorological station at Vatín. Monthly data were calculated as well as evapotranspiration using Pennmann quotation (Barančíková 2005; Barančíková et al. 2014). Simulation of soil organic carbon stock was calculated for two climatic scenarios: M2 - rapid rate of temperature increasing, M3 - less rapid increasing of temperature. Source of climatic scenarios (2000 to 2100) are up- to-data from two global circulation models HadGEM2 and MRI-CGCM3 selected from CMIP5 ensemble (TAYLOR et al. 2012) These projections were prepared using M&Rfi weather generator (used e.g. within RÖTTER et al. 2011) in connection with the Representative Concentration Pathway (RCP) 8.5 greenhouse gas concentration trajectory. Soil data were collected twice a year (spring and autumn) during the period 1999-2015 and calculated according to Falloon et al. (2000) and Falloon & SMITH (2002) as follows:

Initial SOC stock =  $SOC \times BD \times SD$ 

where:

SOC – soil organic carbon content (%)

BD – bulk density (g/cm<sup>3</sup>)

SD - soil depth (cm)

The initial SOC content was used for running Roth C model to equilibrium (10 000 years) under constant environmental conditions. Than the carbon inputs were fitted to match the initial SOC stock, DMP, RMP, BIO, and HUM with different decomposition rate. Organic carbon inputs of plant residues or farmyard manure were calculated according to BIELEK and JURČOVÁ (2010). Data of carbon and radiocarbon ages were received in equilibrium mode (initial soil state, initial radiocarbon age), and were applied to run model in short term mode (1999–2015), and for prediction in long term mode (2015–2100). Total differences between simulated and measured data were calculated according to Loague and Green (1991) as a root mean square error (RMSE).

### RESULTS AND DISCUSSION

Studied Eutric Cambisol was loamy-sand textured, with acid soil reaction, low cation exchange capacity, and low soil colloidal complex saturation. Average measured values of SOC during field experiment 2006–2016 are showed in Figure 1. Humus content was satisfactory but its quality was low, with prevalence of fulvic acids ( $C_{\rm HA}/C_{\rm FA}<1$ ). Humification degree was less than 25%. Soil contains no carbonates. Comparison of soil properties under both studying cropping systems (organic and intensive) is showed in Table 1 and 2. Organic crop sequence was represented by 33.4% of cereals, 16.6% of root crops, 16.6% of

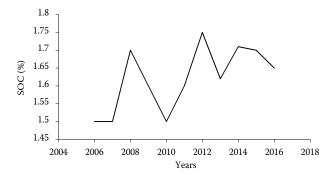


Figure 1. Measured average carbon content in 0-0.20~m during 2006-2016

SOC – soil organic carbon content

Table 1. Basic physical and chemical properties of Eutric Cambisol

Eutric	pH*		CEC	V	Clay content	
Cambisol	$H_2O$	KCl	(cmol/100g)	(%)		
OCS	5.5	4.5	14.2	63.4	22.2	
ICS	5.3	4.3	15	63.3	22	

\*pH/H<sub>2</sub>O – active soil reaction, pH/KCl – exchangeable soil reaction; CEC – cation exchange capacity; V – saturation of soil colloidal complex; OCS – organic crop sequence; ICS – intensive crop sequence

Table 2. Average content of total organic carbon and fractional composition of humic substances in Eutric Cambisol

Eutric	TOC	ΣHS	ΣΗΑ	ΣFA HA/FA		HD
Cambisol	(%)		(g/kg)		ПА/ГА	(%)
OCS	2	4.5	2	2.5	0.8	22.5
ICS	1.8	4.5	2	2.5	0.8	25

TOC – total organic carbon; HS – humic substances; HA – humic acids; FA – fulvic acids; HD – humification degree; OCS – organic crop sequence; ICS – intensive crop sequence

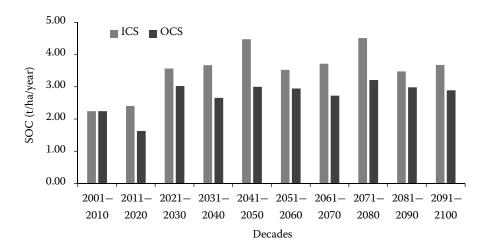


Figure 2. Projected average carbon input (t/ha/year) 0–0.20 m under different management scenario for decades

SOC – soil organic carbon content; OCS – organic crop sequence; ICS – intensive crop sequence

technical crops, and 33.4% of fodder. Intensive crops sequence was represented by 50% of cereals, 16.6% of root crops, and 33.4% of technical crops. Results showed slightly higher quality of humus and soil colloidal complex saturation, and less acidity after ten years of organic farming. Both of them differ in the overall amount of postharvest remains and straw passing every year into the soil. Higher input of plant residues was under ICS management and therefore projected total organic carbon is higher -Figure 2. Typical average yield of grown plants during the selected period is listed in Figure 3. As quoted TESAŘOVÁ et al. (2006) sum of stubble straw and root residues passing every year into the soil at this locality has reached 5.6-3.97 t/ha for winter wheat and spring barley. The postharvest residua of both crops involved 20–30% of the roots. No relationship was found between the total amount of postharvest residua and yields. Root remains of cereals were decomposed under field conditions substantially more slowly than the straw. The decomposition rate was higher under organic farming system. Validation of RothC-26.3 model was done using data from organic farming and calculated RMSE (mean quadratic standard deviation) was 14.90%. Literature data for long-term field experiments are between 2–30% (SMITH *et al.* 1997, 2005, 2007; FALLOON & SMITH 2002; BARANČÍKOVÁ *et al.* 2014). Measured data of soil organic carbon content under organic farming are in good accordance with simulated data.

As it was mentioned before for projection of soil organic carbon content during the period of 2000–2100 we used data from HadGEM2 and MRI-CGCM3

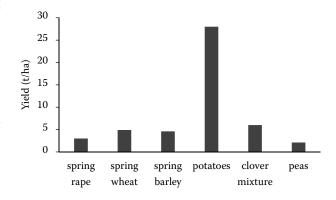


Figure 3. Typical average yield of grown plants at studied locality (intensive crop sequence; t/ha)

Table 3. Temperature simulating over the decades for two climatic scenarios in connection with RPC 8.5

Decades	T (°C)			
Decades	HadGEM2	MRI-CGCM3		
2001–2010	7.54	7.54		
2011-2020	8.30	8.08		
2021-2030	8.43	7.93		
2031-2040	8.87	8.18		
2041-2050	9.86	8.95		
2051-2060	10.20	9.05		
2061-2070	11.06	9.66		
2071-2080	11.90	10.26		
2081-2090	12.38	10.48		
2091-2100	13.34	11.21		

HadGEM2 – hot-dry scenario; MRI-CGCM3 – warm-dry scenario

models – Table 3. Model's input data are listed in Table 4. Simulated prognosis of carbon stock in short term mode (1999–2015) and long term mode

(2015–2100) indicated that at the beginning of the simulated period simulation soil organic carbon content was decreasing. Later (after 2020) higher SOC stock under intensive farming was obtained -Figure 4. Accumulation ability of Eutric Cambisol was evaluated as high and confirmed that type of land management is an important factor influencing soil organic carbon stock (Figures 4 and 5). It should be also stressed that besides crop management and climatic scenario plant input and microbial activity are very important factors as well. In Figure 4 it is showed simulated amount of SOC stock for decades under organic and intensive farming systems (OCS-M2 - organic crop sequence, hot-dry with rapid temperature changes, OCS-M3 - organic crop sequence, warm-dry with less rapid temperature changes, ICS-M2 - intensive crop sequence, hot-dry with rapid temperature changes, ICS-M3 - intensive crop system, warm-dry with less rapid temperature increasing). We can conclude that different farming systems on the same soil type lead to a completely different soil organic carbon stocks. In our case,

Table 4. Development of carbon input under different agronomic scenario over the decades (in t/ha/year)

Decades —		ICS			OCS		
	carbon of PlantRes	carbon of FYM	sum	carbon of PlantRes	carbon of FYM	sum	
2001–2010	1.34	0.91	2.25	1.34	0.91	2.25	
2011-2020	1.95	0.45	2.4	1.63	0	1.63	
2021-2030	2.43	1.13	3.57	1.94	1.08	3.02	
2031-2040	2.42	1.25	3.67	2.01	0.65	2.66	
2041-2050	2.89	1.59	4.48	1.84	1.16	3.00	
2051-2060	2.39	1.13	3.52	1.85	1.08	2.95	
2061-2070	2.47	1.25	3.72	2.07	0.65	2.72	
2071-2080	2.92	1.59	4.51	2.07	1.14	3.21	
2081-2090	2.34	1.13	3.48	1.90	1.08	2.98	
2091-2100	2.43	1.25	3.68	2.24	0.65	2.89	

ICS - intensive crops sequence; OCS - organic crop sequence; PlantRes - plant residue; FYM - fytomass

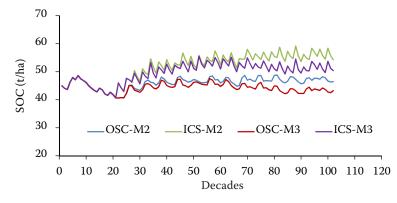


Figure 4. Projected development of soil organic carbon content (SOC) stock in 0–0.20 m for the period 1991–2100

OCS-M2 – organic crop sequence with hot-dry climatic scenario; OCS-M3 – organic crop sequence with war-dry climatic scenario; ICS-M2 – intensive crop sequence with hot dry climatic scenario; ICS-M3 – intensive crop system with warm-dry climatic scenario

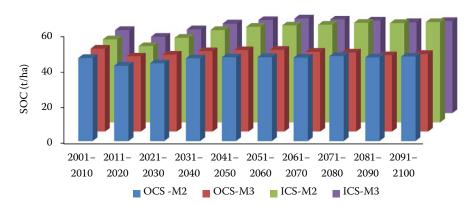


Figure 5. Average carbon stock (t/ha) in 0–0.20 m under different management and climatic scenario for decades OCS-M2 – organic crop sequence and rapid temperature changes; OCS-M3 – organic crop sequence and less rapid temperature changes; ICS-M2 – intensive crop sequence and rapid temperature changes; ICS-M3 – intensive crop system and less rapid temperature increasing; SOC – soil organic carbon content

intensive farming because of higher plant residues input and lower mineralization rate was presented by higher soil organic carbon stock. Organic farming showed higher mineralization rate, and lower organic carbon stock during the projected period. Correlation coefficient between SOC stock and temperature (HadGEM2; hot-dry climatic scenario) was 0.65 in organic farming system. ICS at the same climatic scenario had correlation coefficient 0.76. Similar results were received for the MRI-CGCM3 (warm-dry) climatic scenario. Correlation coefficient R = 0.76 was reached for intensive farming and R = 0.71 for organic farming. Obtained results also confirmed that despite of lower carbon stock in soil under organic farming this management is less influence by climatic conditions to compare with intensive farming system.

# CONCLUSION

Carbon sequestration in soil is an effective strategy to mitigate global climate change. High accumulation potential of carbon in Eutric Cambisol was determined. In spite of less carbon input organic farming was more stable to compare with intensive farming. Intensive farming system was much more effected by climatic condition and plant residues input. Application of RothC-26.3 is a useful tool for carbon stock projection in the long- and short-term mode.

#### References

Baldock J.A., Skjemstad J.O. (1999): Soil organic carbon/soil organic matter. In: Peverill K.I., Sparow L.A., Reuter D.J.

(eds.): Soil Analysis: An Interpretation Manual. Collingwood, CSIRO Publishing: 159–170.

Barančíková G. (2005): Final Report of the Fellowship of Dr. G. Barancikova within the Framework of the MET-AGE Project (EV/10/14A), Louvain-la-Neuve, 2005: 9.

Barančíková G., Skalský R., Koco Š., Halas J., Tarasovičová Z., Nováková M. (2014): Farm-level modelling of soil organic carbon sequestration under climate and land use change. In: Halldórsson G., Bampa F., Porsteinsdóttir A.B. (eds.): Soil Carbon Sequestration for Climate Food Security and Ecosystem Services. Luxembourg, Publications Office of the European Union: 94–100.

Bielek P., Jurčová O. (2010): Methodology for Organic Carbon Balance and Setting up the Organic Fertilization Rates for Agricultural Soils. Bratislava, Soil Conservation and Research Institute. (in Slovak)

Coleman K., Jenkinson D.S. (2005): ROTHC-26.3. A Model for the Turnover of Carbon in Soil. Model Description and Windows Users' Guide, November 1999 Issue (modified April, 2005). Available at http://www.rothamsted.bbsrc.ac.uk/aen/carbon/mod26\_3\_win.pdf (accessed Jan 2018).

Coleman K., Jenkinson D.S., Crocker G.J., Grace P.R., Klir J., Körschens M., Poulton P.R., Richter D.D. (1997): Simulating trends in soil organic carbon in long-term experiments using RothC-26.3. Geoderma, 81: 29–44.

De Liu L., O'Leary G.J., Ma Y., Cowie A., Li F.Y., Mc-Caskill M., Conyer M., Dalal R., Robertson F., Dougherty W. (2016): Modelling soil organic carbon 2. Changes under a range of cropping and grazing farming systems in eastern Australia. Geoderma, 265: 164–175.

Falloon P., Smith P. (2002): Simulating SOC changes in long-term experiments with RothC and century: model evaluation for regional scale application. Journal of Soil Use and Management, 18: 101–111.

- Falloon P., Smith P., Coleman K., Marshall S. (1998): Estimating the size of the inert organic matter pool from total soil organic carbon content for use in the Rothamsted carbon model. Journal of Biology and Biochemistry, 30: 1207–1211.
- Falloon P., Smith P., Coleman K., Marshall S. (2000): How important is inert organic matter for predictive soil carbon modelling using the Rothamsted carbon mode? Journal of Biology and Biochemistry, 32: 433–436.
- IUSS Working Group WRB (2015): World Reference Base for Soil Resources 2014, Update 2015, Rome, FAO. Available at http://www.fao.org/3/a-i3794e.pdf (accessed Jan 2018)
- Jahn R., Blume H.P., Asio V.B., Spaargaren O., Schad P. (2006): Guidelines for Soil Description, 4<sup>th</sup> Ed. Rome, FAO.
- Jenkinson D.S., Hart P.B.S., Rayner J.H., Parry L.C. (1987): Modelling the turnover of organic matter in long-term experiments at Rothamsted. INTELCOL Bulletin, 15: 1–8.
- Jenkinson D.S., Meredith J., Kinyamario J.I., Waren G.P., Wong M.T.H., Harkness D.D., Bol R., Coleman K. (1999): Estimating net primary production from measurements made on soil organic matter. Ecology, 80: 2762–2773.
- Kaczynski R., Siebiele G., Galazka R., Niedzwieck J.M., Polakova S. (2013): Assessment of Soil Organic Carbon Status and Changes in Soils of Polish-Czech Borderland. Brno, Central Institute for Supervising and Testing in Agriculture. (in Czech)
- Keryn P., Polglase P. (2004): Calibration of the RothC model to turnover of soil carbon under eucalypts and pines. In: 3<sup>rd</sup> Australian New Zealand Conf., Dec 5–9, 2004: 1–3. Available at www.regional.org.au/au/assi/ (accessed Jan 2018)
- King J.A., Bradley R.I., Harrison R. (2005): Current trends of soil organic carbon in English arable soils. Journal of Soil Use and Management, 21: 189–195.
- Kučerík J., Šmejkalová D., Čechlovská H., Pekař M. (2007): New insights into aggregation and conformational behaviour of humic substances: Application of high resolution ultrasonic spectroscopy. Organic Geochemistry, 38: 2098–2110.
- Lamar R., deTourdonnet S., Barz P., During R.A., Frielinghaus M., Kolli R., Kubát J., Medveděv V., Netland J., Picard D. (2006): Prospect for conservation agriculture in northern and eastern European countries. Lesson of KASSA. Bibliotheca Fragmenta Agronomica, 11: 77–88.
- Loague K., Green E.E. (1991): Statistical and graphic methods for evaluating solute transport models. Overview and application. Journal Contaminant Hydrology, 7: 51–73.
- Lorenz K., Lal R. (2005): The depth distribution of soil organic carbon in relation to land use and management and the potential of carbon sequestration in subsoil horizons. Advances of Agronomy, 88: 35–66.

- Machmuller M.B., Kramer M.G., Cyle T.K., Hill N., Hancock D., Thomson A. (2015): Emerging land use practices rapidly increase soil organic matter. Nature Communications, 6: 6995.
- Nelson D.W., Sommers L.E. (1996): Total carbon, organic carbon, and organic matter. In: Sparks D.L. *et al.* (eds.): Methods of Soil Analysis. Part 3. Chemical Methods. Madison, SSSA Book Series No. 5, SSSA and ASA: 961–1010.
- Němeček J. *et al.* (2011): Taxonomic Soil System of the Czech Republic. 2<sup>nd</sup> Ed. Prague, Czech Agricultural University Prague. (in Czech)
- Pohanková E., Hlavinka P., Takáč J., Žalud Z., Trnka M. (2015): Calibration and validation of the crop growth model DAISY for spring barley in the Czech Republic. Acta Universitatis Agriculturae et Silviculturae Mendelianae Brunensis, 63: 1177–1186.
- Pospíšilová L., Vlček V., Hybler V., Hábová M., Jandák J. (2016): Standard analytical methods and evaluation criteria of soil physical, agrochemical, biological and hygienic parameters. Folia Universitatis Agriculturae at Silviculturae Mendelianae Brunensis, IX, 2016, 3.
- Rötter R.P., Palosuo T., Pirttioja N.K., Dubrovsky M., Salo T., Fronzek S., Aikasalo R., Trnka M., Ristolainen A., Carter T.R. (2011): What would happen to barely production in Finland if global warming exceeded 4°C? A model-based assessment. European Journal of Agronomy, 35: 205–214.
- Smith P., Smith J.U., Powlson D.S., McGill W.B., Arah J.R.M., Chertov O.G., Coleman K., Franko U., Frolking S., Jenkinson D.S., Jensen L.S., Kelly R.H.M., Klein-Gunnewiek R., Komarov A.S., Li C., Molina J.A.E., Mueller T., Parton W.J., Thorney J.H.M., Whitmore A.P. (1997): A comparison of the performance of nine soil organic matter models using datasets from seven long-term experiments. Geoderma, 81: 153–225.
- Smith J., Smith P., Wattenbach M., Zaehle S., Hiederer R., Jones R.J.A., Montanarella L., Rounsevell M., Reginster I., Ewert F. (2005): Projected changes in mineral soil carbon of European croplands and grasslands, 1990–2080. Global Change Biology, 11: 2141–2152.
- Smith J., Smith P., Wattenbach M., Gottschalk P., Romanenkov V.A., Sevcova L.K., Sirotenko O.D., Rukhovic D.I., Korolova P.V., Romanenko I.A., Lisovoj N.V. (2007): Projected changes in the organic carbon stocks of cropland mineral soils of European Russia and the Ukraine 1990–2070. Global Change Biology, 13: 342–354.
- Smith P., Faloon P., Kutsch W.L. (2010): The role of soils in the Kyoto protocol. In: Kutsch W.L., Bahn M., Heinemeyer A. (eds.): Soil Carbon Dynamics. An Integrated Methodology. Cambridge, Cambridge University Press: 245–256.

Song X., Liu S., Liu Q., Zhang W., Hu Ch. (2014): Carbon sequestration in soil humic substances under long-term 399 fertilization in a wheat-maize system from North China. Journal of Integrative Agriculture, 13: 562–569.

Stevenson F.J. (1994): Humus Chemistry. 2<sup>nd</sup> Ed. Genesis, Composition and Reactions. New York, Wiley.

Taylor K.E., Stouffer R.J., Meehl G.A. (2012): An overview of CMIP5 and the experiment design. Bulletin American Meteorological Society, 93: 485–498

Tesařová M., Kudlička P., Pospíšilová L., Kalhotka L., Hrabě F. (2006): Comparison of mineralisation and humification of postharvest residues of cereals in conventional and organic

cropping practices. Acta Universitatis Agriculturae et Silviculturae Mendelianae Brunensis, 54: 121–126.

Viscarra Rossel R.A., Brus D.J., Lobsey C., Shis Z., McLachlan G. (2016): Baseline estimates of soil organic carbon by proximal sensing: Comparing design-based, model-assisted and model-based inference. Geoderma, 265: 152–163. Zbíral J., Honsa I. *et al.* (2010): Unified Working Instructions. Soil Analysis I. 3<sup>rd</sup> Ed. Brno, Central Institute for Supervising and Testing in Agriculture. (in Czech)

Received for publication April 9, 2018 Accepted after corrections October 8, 2018 Published online February 25, 2019